










## Article

# Depth-Wise Assessment of Soil Fertility and Organic Carbon Under Different Land Use Systems: Implications for Climate Change Adaptation and Resilience in Smallholder Agroecosystems

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## Abstract

This study investigates the influence of various land use systems (LUSs) on soil physico-chemical properties, nutrient dynamics, and soil organic carbon (SOC) stocks in the Central Plain Zone of Uttar Pradesh, India. Soil samples were collected from six distinct LUSs, i.e., fallow, crop-based, horticulture-based, forest-based, vegetable-based, and barren land, and analyzed across three depth intervals (0–15 cm, 15–30 cm, and 30–60 cm). Soil pH increased steadily with depth, ranging from 7.43 to 8.58 at the surface layer to 7.55 to 10.32 in deeper layers. Horticulture-based LUSs recorded the lowest pH, while barren lands had the highest. Electrical conductivity (EC) also rose with depth, ranging from 0.12 to 3.63 dS m<sup>-1</sup>, from the surface to subsoil layers, all below critical salinity thresholds. Soil organic carbon (SOC) content decreased with increasing

soil depth across all land use systems. Among the studied systems, horticulture-based land use recorded the highest SOC content (0.77%), whereas barren land showed the lowest SOC content (0.21%). Due to greater organic matter inputs and reduced disturbances, horticultural systems also exhibited significantly higher levels of macronutrients (N: 17.98 kg ha<sup>-1</sup>, P: 330.45 kg ha<sup>-1</sup>, K: 374.81 kg ha<sup>-1</sup>, S: 84.33 mg ha<sup>-1</sup>) and micronutrients (Fe: 164.12 mg ha<sup>-1</sup>, Mn: 60.89 mg ha<sup>-1</sup>, Cu: 2.85 mg ha<sup>-1</sup>, Zn: 1.80 mg ha<sup>-1</sup>). Bulk density increased slightly with depth (1.46–1.63 Mg m<sup>-3</sup>), while soil moisture content remained relatively stable (43.43% to 42.31%), with moderate variability (CV: 24–27%). The mean total SOC stock was 10.77 t C ha<sup>-1</sup>, ranging from 5.44 to 14.46 t C ha<sup>-1</sup>. Microbial properties also varied among land uses: dehydrogenase activity (DEA), an indicator of microbial functionality, peaked in vegetable-based systems (30.54 µg TPF g<sup>-1</sup>), whereas microbial biomass carbon (MBC) was highest in forest-based systems (184.83 µg g<sup>-1</sup>). Correlation and regression analyses revealed a strong positive relationship between SOC and nutrient availability, with the highest correlation observed for Zn ( $R^2 = 0.99$ ), followed by N ( $R^2 = 0.83$ ) and K ( $R^2 = 0.75$ ). Overall, barren lands showed the poorest soil quality indicators, while horticulture-based systems consistently demonstrated superior soil fertility and carbon sequestration potential. These findings emphasize the critical role of land use management in regulating soil fertility, SOC dynamics, and the long-term sustainability of agro-ecosystems in the region.

**Keywords:** organic carbon storage; land use systems; soil quality; carbon sequestration; soil health; microbial properties

## 1. Introduction

Land use patterns reflect the complex interactions between human activities and the natural environment, highlighting how resources such as forests are exploited over time and across regions [1]. Understanding shifts in land use is critical for developing strategies that harmonize human needs with ecological sustainability [2]. Conversion of forests into agricultural, horticultural, or other land types is a major driver of resource degradation, disrupting nutrient and carbon cycles, reducing soil fertility, and diminishing biodiversity [3]. These impacts are especially pronounced in tropical regions, where climatic conditions and inadequate soil management accelerate organic matter decomposition, nutrient loss, and soil degradation [4]. In India's dry tropical zones, including the Central Plain Zone of Uttar Pradesh, such land use changes compromise soil structure and fertility, threatening crop productivity, food security, and the resilience of socio-ecological systems [5]. Overexploitation of forest products, intensive farming, and livestock grazing further simplify ecosystems and exacerbate degradation. While local socio-economic and cultural factors influence land-use decisions, environmental variables like climate, soil health, and biodiversity remain central drivers [6].

Alterations in land use also have broader ecological consequences, including climate change and biodiversity loss [7]. Soil organic carbon (SOC), a key component of the global carbon cycle, is particularly sensitive to these changes [8]. Maintaining and enhancing SOC is essential for soil fertility, ecosystem functioning, and climate mitigation [8]. Soils act as major terrestrial carbon reservoirs, both absorbing and releasing CO<sub>2</sub> [9]. Initiatives such as the global "4 per 1000" program highlight that even modest increases in SOC can significantly offset greenhouse gas emissions while improving crop yields, emphasizing the importance of sustainable soil management [10,11].

Worldwide, land degradation and agricultural expansion have substantially reduced SOC stocks and contributed to greenhouse gas emissions [12,13]. Adoption of sustainable practices such as conservation tillage, agroforestry, and organic amendments can improve SOC sequestration and soil quality [14,15]. Historical land use changes have substantially reduced global SOC stocks due to imbalances between carbon inputs and mineralization losses [16,17]. Intensive land use generally depletes SOC, whereas restoration and sustainable management practices can improve carbon sequestration and soil resilience, particularly in vulnerable regions such as the Himalayas [18,19].

Soil fertility is closely linked with SOC and plays a critical role in sustainable agriculture and food security [20,21]. Monitoring changes in soil properties under different land-use systems is therefore essential for sustainable land management and productivity assessment [22–24].

In Uttar Pradesh's Central Plain Zone, an agriculturally significant region, land use systems heavily influence soil fertility and crop productivity [25,26]. Practices like intensive plowing, fertilizer application, and crop residue removal further modify soil physico-chemical properties. Despite this, few studies have systematically quantified soil fertility depletion across different land uses in the area. Evaluating soil fertility under varying land use systems is therefore essential for assessing ecosystem health, maintaining long-term food security, and guiding sustainable land management strategies [27]. The novelty of this study lies in its depth-wise evaluation of soil fertility, soil organic carbon fractions, and microbial activity under multiple land use systems in the Central Plain Zone of Uttar Pradesh. The study integrates physical, chemical, biological, and statistical analyses to better understand land use effects on soil health and carbon dynamics.

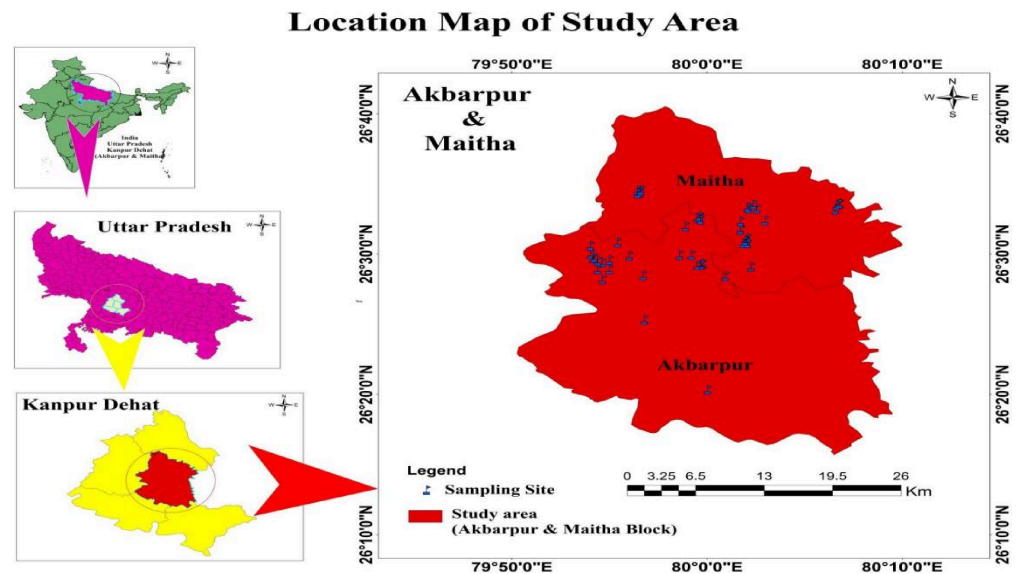
Advanced statistical tools, including PCA, correlation analysis, and heatmap visualization, were used to evaluate depth-wise variations in soil fertility indicators and organic carbon fractions under different land use systems.

## 2. Materials and Methods

### 2.1. Region, Climate, and Soil

The study was conducted in Kanpur Dehat district, which lies within the Upper Ganga Plain (West)/Middle Ganga Plain (West) region. The district is bordered by Kanpur Nagar to the east, Kannauj to the north, Auraiya to the west, and Jalaun and Hamirpur to the south. It covers a total geographical area of 3021 km<sup>2</sup> and lies at an average elevation of 126 m above mean sea level. Geographically, it extends between 26°31' to 35°75.94" N latitude and 79°49" to 84°46.9" E longitude, with the topography sloping from north to southeast, situating it entirely within the lower Doab of the Ganga and Yamuna rivers. The climate of the district is characterized by scorching summers with hot, dry winds, except during the southwest monsoon season. The maximum summer temperature may reach 47 °C, while the minimum is around 30 °C. The average annual rainfall is about 950 mm. The region has an ustic soil moisture regime with a hyperthermic temperature regime. The major soil groups in the area are Haplustalfs, Haplustepts, and Ustorthents, with soil textures ranging from loamy sand to sandy clay loam, which makes it prone to soil–water erosion.

For the study, ten distinct sites were selected from each village, and the district was divided into two blocks—Akbarpur (villages: Muridpur, Gutaha, Tigayi, Bewan, Bhatauli, Raniya, Rura, Lalpur, Raipur, Patari) and Maitha (villages: Kashipur, Aungi, Hathika, Shivali, Sambharpur, Bairi, Bhagpur, Maitha, Devipur, Jugrajpur) (Figure 1). The investigation focused on six major land use systems: fallow land, cultivated crop land, horticulture, forestry, vegetable crop land, and barren land.



**Figure 1.** Locations of soil sample collection.

## 2.2. Soil Sampling and Analysis

For this study, six land use systems (LUSs) were selected from the Central Plain Zone of Kanpur, India: fallow land (LUS1, no crop), crop-based land (LUS2, rice–wheat system), horticultural land (LUS3, mango orchards), forestry land (LUS4, eucalyptus plantations), vegetable-based land (LUS5), and barren land (LUS6). From each LUS, ten sites were chosen, resulting in a total of sixty sampling sites. At each site, three soil pits were arranged in a triangular pattern to reduce spatial variability and ensure representative sampling. At each sampling site, three replicate soil samples were collected from each depth interval and composited into a single representative sample to minimize spatial variability and improve sampling reliability. Soil samples were collected at three depth intervals—0–15 cm (D1), 15–30 cm (D2), and 30–60 cm (D3)—using a khurpi and a stainless steel auger. The three replicate samples from each depth were composited into a single sample per site. Additionally, undisturbed soil cores were obtained with a core sampler from the center of each triangular plot to measure bulk density. Soil moisture content (SMC) was determined using the gravimetric oven-drying method. Fresh soil samples were weighed, oven-dried at 105 °C for 24 h, and reweighed. Soil moisture content was calculated as the percentage loss in weight after drying [27].

Soil samples were air-dried at room temperature, gently crushed to break soil aggregates, and passed through a 2 mm sieve. All subsequent physico-chemical and SOC analyses were conducted on the fine earth fraction (<2 mm) following standard soil analytical procedures. Although the sampled alluvial soils contained negligible visible coarse fragments, the coarse fragment fraction (>2 mm) was not quantified separately. Therefore, SOC stock estimates should be interpreted with caution where coarse fragments may occur. Bulk density was determined using undisturbed core samples after oven drying at 105 °C. Since coarse fragments were negligible in the studied soils, bulk density calculations were based primarily on the fine earth fraction (<2 mm). We also acknowledge that explicit coarse fragment correction using displacement methods would improve SOC stock accuracy in gravel-rich soils. The physical properties analyzed included soil particle size distribution, bulk density, water-holding capacity, and porosity. The chemical properties measured were soil pH, electrical conductivity, organic carbon, available nitrogen, phosphorus, potassium, sulfur, and diethylenetriaminepentaacetic acid (DTPA)-extractable micronutrients (iron, manganese, zinc, and copper). All laboratory analyses were performed using the fine earth fraction (<2 mm) in accordance with standard soil analytical protocols.

### 2.3. Laboratory Analysis of Soil

The Soil Science and Agricultural Chemistry laboratory at C.S. Azad University of Agriculture and Technology, Kanpur 208017, Uttar Pradesh, India, is where the soil samples were examined. At a soil:water ratio of 1:2.5, the pH and electrical conductivity ( $\text{dS m}^{-1}$ ) of the soil were measured using digital pH and EC meters, respectively [28]. Using a diphenyl amine indicator, oxidizable soil organic carbon (SOC) was calculated using the quick titration method developed by Walkley and Black [29]. The dry bulk density (BD) of the soil samples was measured by the core sampling method [30]. Particle size analyses of soils were performed using the hydrometer method, as described by Bouyoucos [31]. The determination of hot water-soluble carbon content was performed in accordance with a previously published protocol [32]. The determination of soil carbohydrate carbon was conducted using the phenol-sulfuric acid method [33]. Soil carbohydrate levels were assessed via spectrophotometry following a sequential two-step acid hydrolysis and subsequent purification of the hydrolysate, largely following a previously published methodology [34]. The alkaline permanganate Kjeldahl method was used to analyze the amount of nitrogen that was available [35]. The techniques outlined by [36] were used to analyze the amount of phosphorus that was available. The soil's accessible potassium level was assessed by the method described by [37] using a flame photometer (1 N ammonium acetate extract), while the soil's available sulfur content was analyzed by the turbidimetric method described by [37]. Using AAS, the DTPA technique was used to extract the available Fe, Mn, Zn, and Cu [38].

### 2.4. Analysis of Soil Organic Carbon Fraction and Stocks

The functional pools/fractions of organic carbon (OC) were estimated using a modified Walkley and Black method [29] as described by Chan et al. [39] using 12.0 N, 18.0 N, and 24.0 N of  $\text{H}_2\text{SO}_4$ . The very labile carbon was estimated using organic C oxidizable by 12.0 N  $\text{H}_2\text{SO}_4$ . The labile carbon was estimated by the differences in SOC oxidizable by 18.0 N and that oxidizable by 12.0 N  $\text{H}_2\text{SO}_4$ . The less labile carbon was estimated by the differences in SOC oxidizable under 24.0 N and that oxidizable by 18.0 N  $\text{H}_2\text{SO}_4$ . The non-labile carbon was estimated by the difference between SOC oxidizable by 36.8 N and that oxidizable by 24.0 N  $\text{H}_2\text{SO}_4$ .

Soil organic carbon (SOC) stock was calculated using SOC concentration, bulk density, and soil depth for the fine earth fraction (<2 mm) following standard procedures [40,41]. Since the sampled alluvial soils contained negligible coarse fragments, a coarse fragment correction factor was not applied in the present study.

$$\text{SOC}_s = \frac{\text{OC} \times \text{BD} \times D}{10} \quad (1)$$

where SOC indicates soil organic carbon stocks ( $\text{t C ha}^{-1}$ ), OC indicates soil organic carbon, BD indicates bulk density, and D indicates depth of soil sample.

### 2.5. Statistical Analysis

At the  $p < 0.05$  level of significance, two-way analysis of variance (ANOVA) was utilized to examine the variation among various soil characteristics and soil organic carbon stocks in relation to changes in the land use system and soil depth of sampling. To determine whether there were any notable differences between the soil samples, Duncan's multiple comparison tests were used [42]. Using IBM SPSS Statistics version 26.0, the Pearson correlation coefficient and principal component analysis were used to determine the correlation between various soil parameters under various land use regimes [43].

Prior to ANOVA, data normality and homogeneity of variance were examined to ensure compliance with statistical assumptions.

### 3. Results

#### 3.1. Variation in Physico-Chemical Properties of Soils Across Land Use Types

The data presented in Table 1 show the range, mean, and coefficient of variation (CV %) of various physico-chemical properties of soils under different land use systems, studied across three depth intervals: 0–15 cm, 15–30 cm, and 30–60 cm. The properties analyzed include soil pH, electrical conductivity (EC), organic carbon (OC), bulk density (BD), and particle size distribution (sand, silt, and clay). These parameters are vital for understanding the fertility, structure, and overall quality of soil under varying land use conditions.

**Table 1.** Mean value of physico-chemical properties of different land use systems at different depths of soil samples.

LUS	Depth	pH	EC (dS m <sup>-1</sup> )	Sand %	Silt %	Clay %
LUS1	D1	7.43 d ± 0.08	0.15 b ± 0.05	39.60 a ± 1.12	40.63 a ± 0.68	19.77 a ± 0.52
	D2	7.51 c ± 0.12	0.42 b ± 0.07	34.43 a ± 1.67	46.88 b ± 1.07	18.68 bc ± 0.86
	D3	7.55 c ± 0.07	0.55 b ± 0.12	26.73 a ± 1.11	52.69 a ± 0.64	20.58 a ± 0.58
LUS2	D1	7.71 c ± 0.20	0.31 b ± 0.06	42.45 a ± 2.28	40.93 a ± 0.68	16.20 b ± 1.60
	D2	7.86 bc ± 0.27	0.34 b ± 0.10	36.63 a ± 1.48	46.35 b ± 1.07	17.02 c ± 1.12
	D3	7.94 b ± 0.47	0.56 b ± 0.12	27.65 a ± 1.12	52.36 a ± 1.34	20.00 a ± 0.39
LUS3	D1	7.77 bc ± 0.12	0.11 b ± 0.02	38.38 a ± 0.89	40.49 a ± 0.57	21.13 a ± 0.86
	D2	7.90 bc ± 0.12	0.18 b ± 0.05	34.65 a ± 1.48	45.75 b ± 1.01	19.60 b ± 0.82
	D3	7.81 bc ± 0.09	0.20 c ± 0.05	26.73 a ± 1.20	51.42 a ± 1.22	21.85 a ± 0.53
LUS4	D1	7.82 b ± 0.07	0.19 b ± 0.03	37.11 a ± 1.11	40.71 a ± 0.88	22.18 a ± 0.86
	D2	8.02 b ± 0.15	0.25 b ± 0.04	37.11 a ± 1.11	45.13 b ± 0.85	19.11 bc ± 0.62
	D3	7.76 bc ± 0.12	0.41 bc ± 0.07	27.30 a ± 1.07	49.34 a ± 1.27	23.36 a ± 0.39
LUS5	D1	7.61 bc ± 0.06	0.13 b ± 0.04	42.80 a ± 1.30	40.39 a ± 1.39	16.81 b ± 1.37
	D2	7.73 c ± 0.15	0.15 b ± 0.04	36.33 a ± 1.48	46.05 b ± 1.07	17.62 bc ± 1.12
	D3	7.83 bc ± 0.10	0.23 c ± 0.02	27.35 a ± 1.12	52.05 a ± 1.34	20.60 a ± 0.39
LUS6	D1	8.58 a ± 0.13	3.63 a ± 0.75	34.23 a ± 18.86	46.06 a ± 17.15	19.71 a ± 1.57
	D2	9.13 a ± 0.12	3.68 a ± 0.83	22.51 b ± 6.13	52.05 a ± 9.38	25.44 a ± 3.73
	D3	10.32 a ± 0.10	4.34 a ± 0.52	27.72 a ± 9.99	49.10 a ± 9.60	23.17 a ± 9.62

**Land Use Systems:** LUS1—fallow land use system, LUS2—crop cultivation-based land use system, LUS3—horticulture-based land use system, LUS4—forest-based land use system, LUS5—vegetable-based land use system, and LUS6—barren land use system. **Depth:** D1 (0–15 cm), D2 (15–30 cm), and D3 (30–60 cm). Letters denote significant difference among land use systems ( $p < 0.05$ ).

The findings show that LUS3 (horticulture-based land use system) had the lowest pH, and LUS6 (barren land use system) had the highest pH. However, when compared to other soil fertility indicators across different land uses, soil pH showed the lowest mean value, suggesting the least amount of variation among the soil samples that were gathered. According to the rating chart taken from [44], soils in forest-based land use systems (pH 7.61), crop-based land use systems (pH 7.71), vegetable-based land use systems (pH 7.77), and horticulture-based land use systems (pH 7.82) were classified as slightly alkaline to moderately alkaline. Land use patterns were found to have a substantial effect on soil pH ( $p < 0.05$ ). On the surface (0–15 cm), however, there was no discernible variation

in soil pH between horticulture-based and vegetable-based land use systems. A mildly to severely alkaline character was noted in the subsurface soil. This outcome is consistent with the findings from [44], which concluded that soil pH rose with soil depth, which is linked to decreased weathering rates and greater carbonate levels.

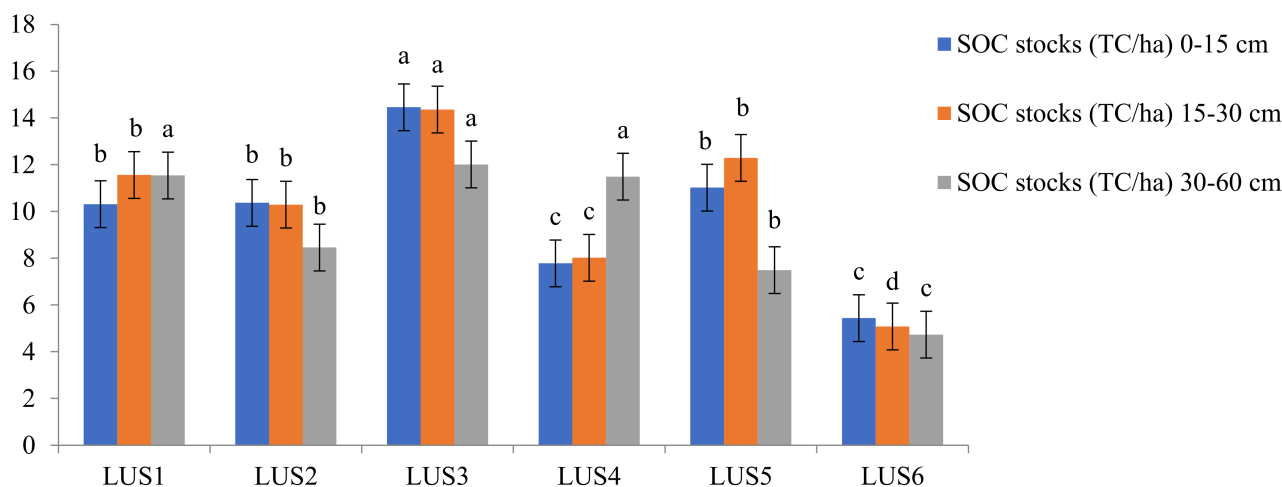
An important measure of soil salinity, soil electrical conductivity (EC), can reveal information on the availability of nutrients in the soil [45]. The recorded soil EC values (measured in  $\text{dS m}^{-1}$ ) in the Central Plain Zone of Uttar Pradesh ranged from 0.12 to 3.63 under the studied LUSs. This suggests that the EC is not the limiting factor for plant growth and crop productivity in the study area, because the concentration of soluble salts was below the critical levels at which the growth and productivity of the majority of the cultivated crops are unaffected [46,47]. The forest-based land use system (LUS4) had the lowest EC value, while the barren land use system (LUS6) had the highest. Results pertaining to EC are shown in Table 1, which shows that no significant differences ( $p < 0.05$ ) were found between the fallow land use system (LUS1) and the crop-based land use system (LUS2), horticulture-based land use system (LUS3), and vegetable-based land use system (LUS5). The presence of salts in the soil, which greatly increases its electrical conductivity, is the cause of the higher EC under the barren land-use regime. An analysis of soil EC ( $\text{dS m}^{-1}$ ) across various land use systems revealed a clear order: LUS2 (0.31) > LUS4 (0.19) > LUS1 (0.15) > LUS5 (0.13). Salts like calcium and sodium chloride dissociate into ions in water, contributing to increased soil conductivity [48]. Across all land uses, this pattern was also seen in the subsurface soil, where EC progressively rose with soil depth [49]. Assessing soil salinity and nutrient dynamics in the research area is made easier with an understanding of these differences in soil EC under various land use systems.

Soil organic carbon content varied significantly among land use systems, with higher SOC under horticultural- and forest-based systems and lower SOC under barren conditions (Figure 2). Greater organic matter input, litter deposition, and reduced soil disturbance in perennial systems likely enhanced SOC accumulation and nutrient retention. The larger amount of SOC in the horticulture land use system among the cultivated LUSs may be caused by fewer disturbances and a higher contribution of plant biomass into the soil in the form of leaf litter residue [50]. Continuous farming without adequate organic matter input might result in decreased organic carbon, which can lead to a decline in soil organic carbon and less vegetation. Soil health can also be negatively impacted by the use of chemical pesticides and fertilizers. These results are consistent with earlier research by [51], who found that forest area in Nepal's Nuwakot and Chitwan districts had higher levels of soil organic matter (OM) than crop land.

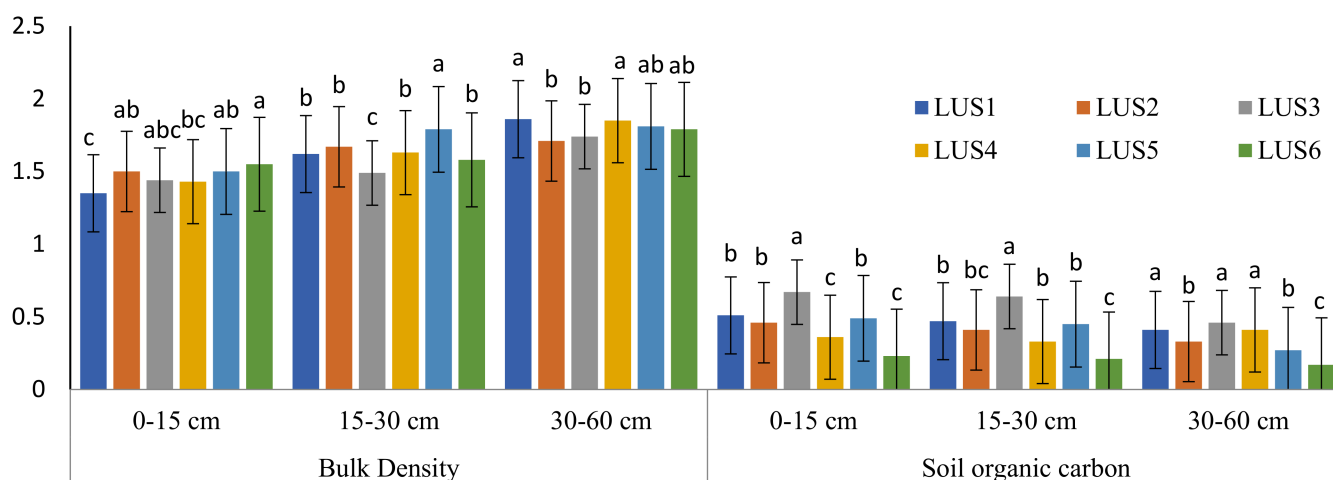
Bulk density (BD) rose from 1.35–1.55  $\text{Mg m}^{-3}$  (mean  $1.46 \pm 0.07$ ) at 0–15 cm to 1.49–1.79  $\text{Mg m}^{-3}$  (mean  $1.63 \pm 0.10$ ) at 15–30 cm and then dipped slightly to 1.71–1.86  $\text{Mg m}^{-3}$  (mean  $1.49 \pm 0.07$ ) at 30–60 cm (Figure 3). The peak in the mid-layer may reflect compaction or lower organic matter, which can impede root growth and water flux. The slight decrease at depth may result from textural or structural changes. Low CVs (4–6%) suggest fairly uniform BD across land uses. Higher BD at the middle depth may be due to compaction and reduced organic matter, whereas a slight reduction at deeper layers could be linked with textural differences and moisture retention. Similar patterns were documented in Indo-Gangetic soils [52].

Particle size distribution (Table 1) shifted toward finer fractions with depth. Sand content decreased from 34.23–42.80% (mean  $39.10 \pm 3.27$ ) at 0–15 cm to 26.73–27.72% (mean  $27.25 \pm 0.43$ ) at 30–60 cm. Conversely, silt rose from 40.39–46.06% (mean  $41.54 \pm 2.22$ ) to 49.10–52.69% (mean  $51.16 \pm 1.56$ ), and clay from 16.20–22.18% (mean  $19.30 \pm 2.36$ ) to 20.00–23.36% (mean  $21.59 \pm 1.43$ ). This indicates textural illuviation, where finer particles

migrate downward and accumulate in subsoil horizons [53]. Such textural stratification directly influences soil water-holding capacity, aeration, and nutrient availability.



**Figure 2.** Soil organic carbon stocks ( $t\ C\ ha^{-1}$ ) in the different land use systems at different depths of soil samples. Lines indicate the standard error of the mean. Bars marked with different letters for a given variable are significantly ( $p < 0.05$ ) different by Tukey’s post hoc test.



**Figure 3.** Soil bulk density ( $Mg\ m^{-3}$ ) and soil organic carbon (%) in the different land use systems at different depths of soil samples. Lines indicate the standard error of the mean. Bars marked with different letters for a given variable are significantly ( $p < 0.05$ ) different by Tukey’s post hoc test.

These variations indicate that vegetation cover, organic matter input, and land use intensity strongly influence soil physico-chemical properties and nutrient dynamics. Horticulture- and forest-based systems generally improved soil quality parameters, whereas barren lands showed signs of degradation and lower soil fertility.

### 3.2. Soil Organic Carbon Fractions and Stock Dynamics in Various Land Use Systems

The data presented in Table 2 show the mean values and variability of different soil organic carbon (SOC) fractions and related soil biological properties across three soil depths (0–15 cm, 15–30 cm, and 30–60 cm) under different land use systems. Soil organic carbon fractions are categorized into very labile, labile, less-labile, and non-labile carbon. Additionally, parameters like hot water-soluble carbon (HWSC), acid-hydrolyzable humic carbon (AHHC), microbial biomass carbon (MBC), dehydrogenase activity (DEA), and soil organic carbon (SOC) stocks were measured to better understand soil health and fertility status.

**Table 2.** Mean values of soil organic carbon fractions under different land use systems at different depths of soil samples.

Depth	LUS	Carbon Fraction (gm kg <sup>-1</sup> )				Hot Water-Soluble Carbon (HWSC)	Acid-Hydrolyzable Humic Carbon (AHHC)	Microbial Biomass Carbon (MBC)	Dehydrogenase Activity (DEA)	Soil Organic Carbon Stocks
		Very Labile	Labile	Less Labile	Non-Labile	mg kg <sup>-1</sup>				
D	LUS1	3.09 b	1.76 c	1.93 b	2.02 b	142.23 b	589.21 b	164.76 b	30.08 a	10.31 b
	LUS2	3.10 ab	1.89 b	1.89 b	2.02 b	137.46 c	561.12 c	172.51 b	26.76 b	10.37 b
	LUS3	3.18 a	2.07 a	2.16 a	2.18 a	149.72 a	604.74 a	182.91 a	24.54 bc	14.46 a
	LUS4	3.07 b	1.88 b	1.89 b	2.18 a	129.60 d	551.96 c	184.83 a	22.60 cd	7.78 bc
	LUS5	3.04 b	1.75 c	1.87 b	2.15 a	149.57 a	579.95 b	182.18 a	30.54 a	11.02 b
	LUS6	2.13 c	1.37 d	1.90 b	1.84 c	112.27 e	201.21 d	108.58 c	20.80 d	5.44 c
D2	LUS1	2.96 ab	1.66 b	2.16 b	1.88 b	123.03 b	359.44 a	130.75 a	22.30 ab	11.56 b
	LUS2	2.93 abc	1.67 b	2.16 b	1.88 b	123.40 b	299.84 c	127.06 ab	23.36 ab	10.29 b
	LUS3	3.04 a	1.89 a	2.47 a	2.05 a	129.60 a	335.86 b	128.95 ab	21.56 b	14.36 a
	LUS4	2.84 c	1.68 b	2.16 b	2.05 a	123.03 b	301.48 c	134.15 a	17.96 c	8.02 c
	LUS5	2.90 bc	1.66 b	2.08 c	1.93 b	129.40 a	330.84 b	121.52 b	24.10 a	12.29 b
	LUS6	2.17 d	1.20 c	1.93 d	1.71 c	100.73 c	134.22 d	101.95 c	16.20 c	5.08 d
D3	LUS1	2.50 b	1.45 a	2.41 a	1.75 b	101.83 c	250.38 a	105.00 b	17.38 a	11.54 a
	LUS2	2.50 b	1.45 a	2.48 a	1.75 b	109.72 b	200.00 c	104.16 b	17.26 a	8.46 b
	LUS3	2.70 a	1.70 a	2.49 a	1.91 a	112.28 a	249.08 a	104.72 b	18.42 a	12.01 a
	LUS4	2.61 ab	1.39 a	2.42 a	1.91 a	101.81 c	251.55 a	111.75 a	13.80 b	11.49 a
	LUS5	2.64 a	1.46 a	2.41 a	1.76 b	109.73 b	232.11 b	105.66 b	17.72 a	7.49 b
	LUS6	2.03 c	3.02 a	2.16 b	1.49 c	90.87 d	103.30 d	91.99 c	14.82 b	4.73 c

**Land Use Systems:** LUS1—fallow land use system, LUS2—crop cultivation-based land use system, LUS3—horticulture-based land use system, LUS4—forest-based land use system, LUS5—vegetable-based land use system, and LUS6—barren land use system. **Depths:** D1 (0–15 cm), D2 (15–30 cm), and D3 (30–60 cm). **HWSC**—hot water-soluble carbon, **AHHC**—acid-hydrolyzable humic carbon, **MBC**—microbial biomass carbon, **DEA**—dehydrogenase enzyme activity, and **SOC**—soil organic carbon. **Units:** mg kg<sup>-1</sup>—milligrams per kilogram, µg g<sup>-1</sup>—micrograms per gram. Letters denote significant difference among land use systems ( $p < 0.05$ ).

It was observed that, at 0–15 cm depth, among all carbon fractions, very labile carbon was highest, followed by the non-labile, less labile, and labile fractions. It can be depicted from the data that soil from LUS3 yields the maximum carbon, viz., 3.18 g kg<sup>-1</sup>, 2.07 g kg<sup>-1</sup>, 2.16 g kg<sup>-1</sup>, and 2.18 g kg<sup>-1</sup> of the VL, L, LL, and NL fractions, respectively. In contrast, minimum soil organic carbon, viz., 2.13 g kg<sup>-1</sup>, 1.37 g kg<sup>-1</sup>, and 1.84 g kg<sup>-1</sup> of the VL, L, and NL fractions was noted for LUS6, except for the LL fraction (1.87 g kg<sup>-1</sup>) from LUS5. However, different soil organic carbon fractions in the rest of the LUS were found to be significantly at par with each other. Similar trends were observed in the case of hot water-soluble carbon (HWSC), which varied from 112.27 to 149.72 mg kg<sup>-1</sup> with a mean of 136.81 mg kg<sup>-1</sup>, acid-hydrolyzable humic carbon (AHHC), 514.70 mg kg<sup>-1</sup> (201.21–604.74 mg kg<sup>-1</sup>), and SOC, 10.77 t C ha<sup>-1</sup> (5.44–14.46 t C ha<sup>-1</sup>). However, the microbial biomass carbon (MBC) and dehydrogenase activity (DEA) fractions were recorded as maximum in LUS4 (184.83 µg g<sup>-1</sup>) and LUS5 (30.54 µg TPF g<sup>-1</sup>), respectively, while the minimums of these fractions were found in LUS6 (108.58 µg g<sup>-1</sup> and 20.80 µg TPF g<sup>-1</sup> day<sup>-1</sup>). These findings are consistent with recent studies reporting that labile carbon fractions are highly sensitive indicators of land use changes and soil health [54,55]. The observed trends agree with previous reports indicating that perennial vegetation and reduced soil disturbance enhance SOC stabilization and microbial activity. Similar observations were reported by [56], who found that very labile and labile carbon fractions decrease markedly under continuous cultiva-

tion. In addition, ref. [57] highlighted that agroecosystem management practices play a crucial role in regulating microbial biomass carbon (MBC) and hydrolysable carbon pools. Further studies [58] have also demonstrated that hot water-soluble carbon (HWSC) and acid-hydrolyzable humic carbon (AHHC) are closely linked to soil functional stability and enzymatic activity. Taken together, the greater microbial biomass carbon (MBC) and dehydrogenase activity (DEA) fractions observed in certain land use systems in the present study support the growing evidence that soil microbial activity provides a dependable indicator of carbon cycling across varied land management practices [59]. Higher SOC under perennial land use systems suggests improved carbon sequestration potential and greater resilience against soil degradation under climate variability.

SOC fractions generally decreased with soil depth across all land use systems, whereas perennial systems maintained relatively higher carbon pools than barren lands. However, different soil organic carbon fractions in the remaining LUSs were found to be significantly at par with each other. Hot water-soluble carbon followed the same pattern, showing greater values under horticultural- and vegetable-based systems compared to barren land. However, acid-hydrolyzable humic carbon (AHHC), microbial biomass carbon (MBC), and dehydrogenase activity (DEA) fractions exhibited maximums in LUS1 ( $359.44 \text{ mg kg}^{-1}$ ), LUS4 ( $134.15 \text{ } \mu\text{g g}^{-1}$ ), and LUS5 ( $24.10 \text{ } \mu\text{g TPF g}^{-1}$ ), respectively, while the minimums of these fractions were found in LUS6 ( $134.22 \text{ mg kg}^{-1}$ ,  $101.95 \text{ } \mu\text{g g}^{-1}$ , and  $16.20 \text{ } \mu\text{g TPF g}^{-1}$ , respectively). These findings are in line with recent reports emphasizing the depth-dependent stabilization of soil organic carbon fractions and their sensitivity to land use practices [60]. The higher stability of recalcitrant fractions, such as the LL- and non-labile carbon fractions, in cropland and agroforestry systems compared to degraded lands has also been highlighted in recent studies. Moreover, microbial biomass and enzyme-mediated fractions were reported to be more responsive indicators of soil quality under diversified land use systems [61]. These outcomes collectively suggest that soil organic carbon stock fractions at subsurface layers serve as reliable markers for evaluating the sustainability of land use systems in semi-arid and alluvial plains [54].

At 30–60 cm of depth, a similar trend was noticed with respect to all fractions of carbon in different LUSs. Maximum carbon, viz.,  $2.70 \text{ g kg}^{-1}$ ,  $2.49 \text{ g kg}^{-1}$ , and  $1.91 \text{ g kg}^{-1}$  of the very labile, less labile, and non-labile fractions, respectively, was found in LUS3, except for the labile fraction, which exhibited a maximum ( $3.02 \text{ g kg}^{-1}$ ) in LUS6, while minimum carbon, viz.,  $2.03 \text{ g kg}^{-1}$ ,  $2.16 \text{ g kg}^{-1}$ , and  $1.49 \text{ g kg}^{-1}$  of the very labile, labile, less labile, and non-labile carbon fractions was observed in LUS6 and the labile fraction in LUS4 ( $1.39 \text{ g kg}^{-1}$ ). However, different soil organic carbon fractions in all other LUSs were found to be significantly at par with each other. A similar trend was observed in the case of hot water-soluble carbon (HWSC), which varied from  $90.87$  to  $112.28 \text{ mg kg}^{-1}$  with a mean of  $104.37 \text{ mg kg}^{-1}$  and soil organic carbon stocks of  $9.29 \text{ t C ha}^{-1}$  ( $4.73$ – $12.01 \text{ t C ha}^{-1}$ ). However, the acid-hydrolyzable humic carbon (AHHC), microbial biomass carbon (MBC), and dehydrogenase activity (DEA) fractions ranged from LUS6 ( $103.30 \text{ mg kg}^{-1}$ ) to LUS4 ( $251.55 \text{ mg kg}^{-1}$ ), LUS6 ( $91.99 \text{ } \mu\text{g g}^{-1}$ ) to LUS4 ( $111.75 \text{ } \mu\text{g g}^{-1}$ ), and LUS4 ( $13.80 \text{ } \mu\text{g TPF g}^{-1}$ ) to LUS3 ( $18.42 \text{ } \mu\text{g TPF g}^{-1}$ ), respectively. The observed differences in soil organic carbon fractions and microbial properties across land use systems highlight the important role of vegetation type, residue input, and management practices in regulating soil carbon stabilization and microbial activity. Systems with greater biomass input and reduced disturbance supported higher SOC accumulation and improved biological activity.

A summary of the major soil quality indicators under different land use systems is presented in Table 3 to facilitate comparative interpretation of overall soil health status.

**Table 3.** Summary of major soil quality indicators under different land use systems.

Soil Quality Indicator	Fallow Land (LUS1)	Crop-Based Land (LUS2)	Horticulture-Based Land (LUS3)	Forest-Based Land (LUS4)	Vegetable-Based Land (LUS5)	Barren Land (LUS6)
pH	Moderately Alkaline	Moderately Alkaline	Slightly Alkaline	Slightly Alkaline	Slightly Alkaline	Highly Alkaline
Electrical Conductivity (dS m <sup>-1</sup> )	Low	Moderate	Low	Lowest	Low	Highest
Soil Organic Carbon (%)	Moderate	Moderate	Highest	High	High	Lowest
Bulk Density (Mg m <sup>-3</sup> )	Moderate	Moderate	Lower	Moderate	Moderate	Higher
Available Nitrogen (kg ha <sup>-1</sup> )	Moderate	Moderate	Highest	High	High	Lowest
Available Phosphorus (kg ha <sup>-1</sup> )	Moderate	Moderate	Highest	High	High	Lowest
Available Potassium (kg ha <sup>-1</sup> )	Moderate	Moderate	Highest	High	High	Lowest
Available Sulfur (mg kg <sup>-1</sup> )	Moderate	Moderate	Highest	High	High	Lowest
Micronutrient Availability (Fe, Mn, Zn, Cu)	Moderate	Moderate	Highest	High	High	Lowest
Hot Water-Soluble Carbon (mg kg <sup>-1</sup> )	High	Moderate	Highest	Moderate	Highest	Lowest
Microbial Biomass Carbon (µg g <sup>-1</sup> )	High	Moderate	High	Highest	High	Lowest
Dehydrogenase Activity (µg TPF g <sup>-1</sup> day <sup>-1</sup> )	Highest	Moderate	Moderate	Lower	Highest	Lowest
SOC Stock (t C ha <sup>-1</sup> )	Moderate	Moderate	Highest	Moderate	High	Lowest
Overall Soil Quality Status	Moderate	Moderate	Excellent	Good	Good	Poor

Note: The table summarizes the relative performance of major soil quality indicators under different land use systems based on mean values across soil depths.

### 3.3. Principal Component Analysis

Principal component analysis (PCA) revealed clear multivariate differentiation among samples, with the first two components explaining 81.6% of total variance (Dim1 = 64.3%, Dim2 = 17.3%) (Figure 4). Dim1 primarily captured variability associated with key soil and physiological indicators, effectively separating Cluster 1 from Clusters 2 and 3. Samples in Cluster 1 were positively associated with variables such as the very labile and non-labile carbon fractions, hot water-soluble carbon (HWSC), acid-hydrolyzable humic carbon, MBC, and DEA, indicating stronger relationships with microbial activity and carbon fractions (Figure 5). In contrast, Cluster 2 samples were positioned on the negative side of Dim1, suggesting distinct structural or compositional characteristics. Dim2 further discriminated Cluster 3, driven largely by SOC and CFLC contributions, highlighting differences linked to soil organic carbon dynamics. The clustering pattern demonstrates substantial heterogeneity across treatments or land use systems, reflecting differential responses in biochemical and microbial attributes. Overall, the PCA biplot confirms that the measured variables effectively capture system-level variability and provide robust discrimination among sample groups.

These results demonstrate that multivariate statistical approaches effectively differentiate land use systems based on their soil biochemical and carbon-related properties.

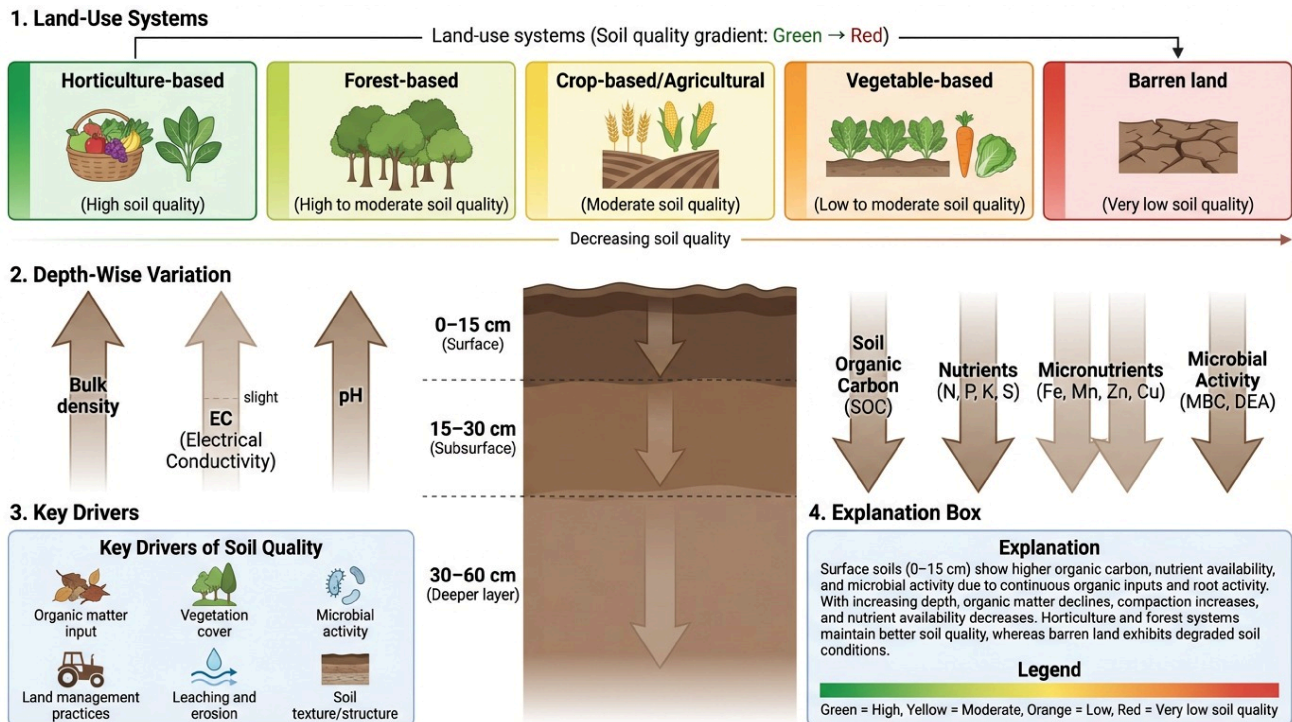


Figure 4. Variations in soil quality indicators, nutrient status, and microbial activity across different land use systems and soil depths.

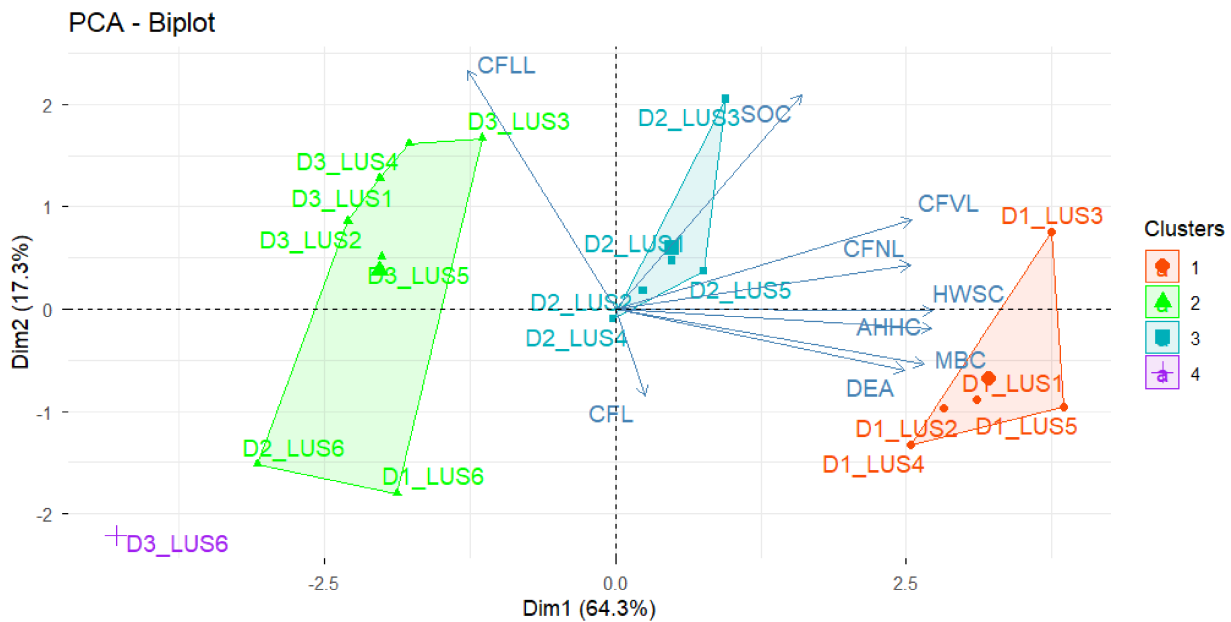
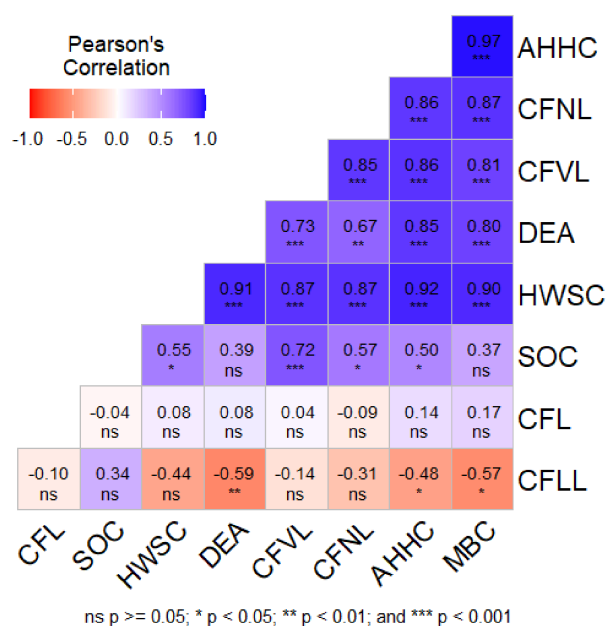


Figure 5. Principal component analysis of organic carbon fractions under different land use systems across varying soil depths. Land Use Systems: LUS1—fallow land use system, LUS2—crop cultivation-based land use system, LUS3—horticulture-based land use system, LUS4—forest-based land use system, LUS5—vegetable-based land use system, and LUS6—barren land use system. Depths: D1 (0–15 cm), D2 (15–30 cm), and D3 (30–60 cm). HWSC—hot water-soluble carbon, AHHC—acid-hydrolyzable humic carbon, MBC—microbial biomass carbon, DEA—dehydrogenase enzyme activity, SOC Stocks—soil organic carbon stocks, CFVL—carbon fraction very labile, CFL—carbon fraction labile, CFLL—carbon fraction less labile, CFNL—carbon fraction Non-labile.

### 3.4. Pearson's Correlation

Pearson's correlation analysis revealed strong and statistically significant relationships among the measured soil biochemical and carbon-related variables (Figure 6). Microbial and enzymatic indicators exhibited particularly high positive correlations. AHHC showed a very strong association with MBC ( $r = 0.97, p < 0.001$ ), indicating a close linkage between hydrolytic activity and microbial biomass. Similarly, HWSC was strongly correlated with DEA ( $r = 0.91, p < 0.001$ ), CFVL ( $r = 0.87, p < 0.001$ ), CFNL ( $r = 0.92, p < 0.001$ ), and MBC ( $r = 0.90, p < 0.001$ ), highlighting the central role of water-soluble carbon in regulating microbial and enzymatic processes. DEA also demonstrated strong positive correlations with CFVL ( $r = 0.85, p < 0.001$ ), CFNL ( $r = 0.86, p < 0.001$ ), AHHC ( $r = 0.85, p < 0.001$ ), and MBC ( $r = 0.80, p < 0.001$ ). SOC displayed moderate but significant positive correlations with CFVL ( $r = 0.72, p < 0.001$ ), CFNL ( $r = 0.57, p < 0.05$ ), and AHHC ( $r = 0.50, p < 0.05$ ), suggesting that increases in organic carbon enhance microbial functionality.



**Figure 6.** Pearson's correlation of organic carbon fractions under different land use systems across varying soil depths. HWSC—hot water-soluble carbon, AHHC—acid-hydrolyzable humic carbon, MBC—microbial biomass carbon, DEA—dehydrogenase enzyme activity, SOC stocks—soil organic carbon stocks, CFVL—carbon fraction very labile, CFL—carbon fraction labile, CFLL—carbon fraction less labile, CFNL—carbon fraction non-labile.

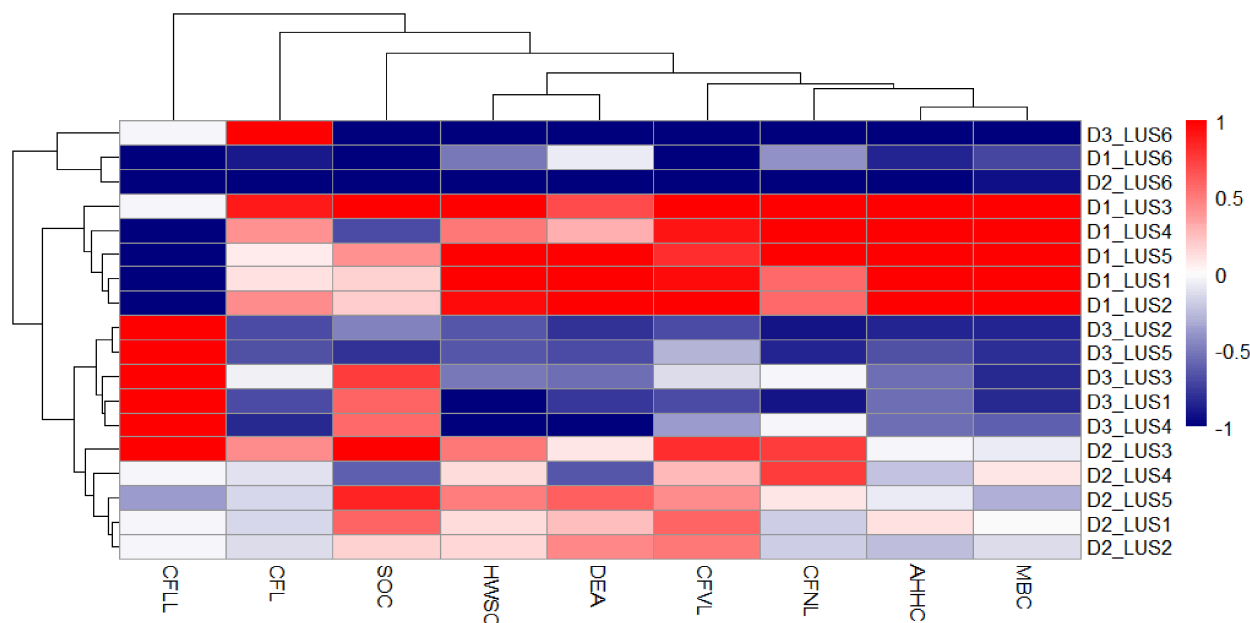
In contrast, CFLL showed significant negative correlations with DEA ( $r = -0.59, p < 0.01$ ), AHHC ( $r = -0.48, p < 0.05$ ), and MBC ( $r = -0.57, p < 0.05$ ), indicating that labile carbon depletion may constrain microbial activity. Overall, the results emphasize strong coupling among carbon fractions, enzyme activity, and microbial biomass.

The strong positive associations among SOC fractions, microbial biomass, and enzyme activity indicate close interactions between soil carbon availability and microbial functioning under different land use systems.

### 3.5. Heatmap Analysis

The heatmap, coupled with hierarchical clustering, illustrates pronounced variability in soil biochemical and carbon fraction attributes across samples. Two major groupings were evident. The first cluster, dominated by D1\_LUS samples (D1\_LUS1-D1\_LUS5), exhibited consistently high standardized values (red spectrum) for HWSC, DEA, CFVL, CFNL, AHHC, and MBC, indicating enhanced microbial biomass and enzymatic activity

(Figure 7). Within this group, D1\_LUS3 and D1\_LUS5 showed particularly strong intensities, reflecting elevated SOC-associated processes.



**Figure 7.** Heatmap analysis of organic carbon fractions under different land use systems across varying soil depths. Land Use Systems: LUS1—fallow land use system, LUS2—crop cultivation-based land use system, LUS3—horticulture-based land use system, LUS4—Forest-based and use system, LUS5—vegetable-based land use system, and LUS6—barren land use system. Depths: D1 (0–15 cm), D2 (15–30 cm), and D3 (30–60 cm). HWSC—hot water-soluble carbon, AHHC—acid-hydrolyzable humic carbon, MBC—microbial biomass carbon, DEA—dehydrogenase enzyme activity, SOC stocks—soil organic carbon stocks, CFVL—carbon fraction very labile, CFL—carbon fraction labile, CFLL—carbon fraction less labile, CFNL—carbon fraction non-labile.

In contrast, a second cluster comprising most D3\_LUS samples (D3\_LUS1–D3\_LUS6) displayed predominantly low standardized scores (blue spectrum), especially for SOC, HWSC, CFNL, AHHC, and MBC, suggesting comparatively reduced microbial and biochemical functioning. Negative deviations approaching  $-1$  were frequent for enzyme-related variables, highlighting potential carbon limitation or slower turnover dynamics.

Samples from D2\_LUS formed an intermediate grouping, characterized by moderate standardized values (white to light red), reflecting transitional system behavior. Notably, SOC and CFLL exhibited heterogeneous patterns, implying differential carbon stabilization mechanisms. Overall, the clustering structure underscores strong co-variation among labile carbon fractions, enzyme activities, and microbial biomass, revealing clear treatment or land-use effects.

Overall, the heatmap analysis confirms substantial variability in soil biochemical properties across land use systems and soil depths, emphasizing the influence of land management practices on soil quality and carbon dynamics.

### 3.6. Nutrient Status of Soil Under Various LUSs

Different LUSs had a direct impact on macronutrient availability, and soil depth affected the amount of accessible N, P, K, and S (Table 4). In the surface soils (0–15 cm) of LUSs under the horticulture-based land use system, available N, P, K, and S were significantly ( $p < 0.05$ ) greater; however, S was recorded under the vegetable-based land use system, and the lowest S content was recorded under the barren land use system of soil. The ranges of available N ( $\text{kg ha}^{-1}$ ) and P ( $\text{kg ha}^{-1}$ ) content were  $17.98 \text{ a} \pm 0.07$ – $12.56 \text{ c} \pm 1.31$

and  $330.45 \text{ a} \pm 7.87$ – $122.79 \text{ d} \pm 11.41$ , respectively. In horticulture-based land use systems, the soil's K ( $\text{kg ha}^{-1}$ ) concentration was highest, whereas in forest-based land use systems, it was lowest. The soil's S ( $\text{Mg ha}^{-1}$ ) content varied from  $84.33 \text{ a} \pm 2.69$  to  $15.15 \text{ c} \pm 1.26$ , with the highest value found in land use systems based on vegetables and the lowest in land use systems based on no vegetation [62]. Because the surface layer (0–15 cm) contains more organic matter, surface soil has more nutrient availability than lower layers, resulting in superior soil quality [63]. The addition of more organic matter through leaf litter, which causes the nutrients in the soil to mineralize, led to a higher content of N, P, and K in the horticulture-based land system [64].

Horticultural land-use soils have significantly higher available N content than crop-cultivated, vegetable-based, forest-based, fallow-land, and barren land use soils. This is because these LUSs have higher organic matter input from the fall of leaf litter and the deposition of root biomass, which has improved the mineralization of nitrogen. In contrast, intensive tillage techniques in agriculture decrease the availability of nitrogen by speeding up the oxidation of soil organic matter [65]. Additionally, a barren land use system with reduced plant cover speeds up the leaching loss of nitrogen through rainfall and irrigation [66]. Under various LUSs, a declining trend in available N with soil depth may be caused by both nitrogen immobilization and a decrease in soil organic carbon at lower soil depths [48]. Both a high amount of soil organic matter, which causes the release of organic phosphorus, and the cumulative effects of continuous phosphoric fertilizer application may be the cause of the horticultural LUS's relatively higher available P [67]. Because of the increased amount of organic matter in the form of leaf litter fall and root biomass deposition, which aids in the release of bound-K and the solubilization of insoluble forms of potassium in the soil, natural forests and horticultural LUSs may have relatively higher available K. Leaching losses and soil degradation may be the cause of the reduced amount of accessible K in the barren land use system [68]. In addition to having decreased potassium content due to high salinity or alkalinity, salts can hinder plants' ability to absorb nutrients, including potassium [69].

Because the amount of organic matter varies with different LUSs, DTPA-extractable micronutrients were impacted by different LUSs at different soil depths. At varying soil depths, the impacts of each LUS on soil micronutrients were significant ( $p < 0.05$ ). Next, compared to the other land uses, the surface layer of the fallow-based land use system, vegetable-based land use system, fallow-based land use system, and crop-based land use system had the significantly highest mean Fe, Mn, Cu, and Zn content ( $164.12 \text{ a} \pm 5.35$ ,  $60.89 \text{ a} \pm 5.43$ ,  $2.85 \text{ a} \pm 0.30$ , and  $1.80 \text{ a} \pm 0.23$ )  $\text{mg kg}^{-1}$ , while the subsurface layers (30–60 cm) of the barren land use system had the lowest mean iron content (Table 2). The available Fe and Mn contents of the soils of the three land uses in the research region fall into the high range, while the available Mn is in the medium range. Assessments of soil-accessible Fe and Mn [70] indicate that the chemical characteristics of the accessible elements Mn and Fe in tropical soils are comparable. In contrast to Fe, the surface layer of the cultivated land had the greatest mean accessible Mn concentration in the research area when compared to the other land use layers. As a result, the available Fe content appears to be more closely associated with the OM content. Due to the frequent input of organic matter in the form of leaf litter and root deposition, which raises microbial activity in the soil and makes micronutrients more readily available, SF soils showed greater levels of micronutrients [67].

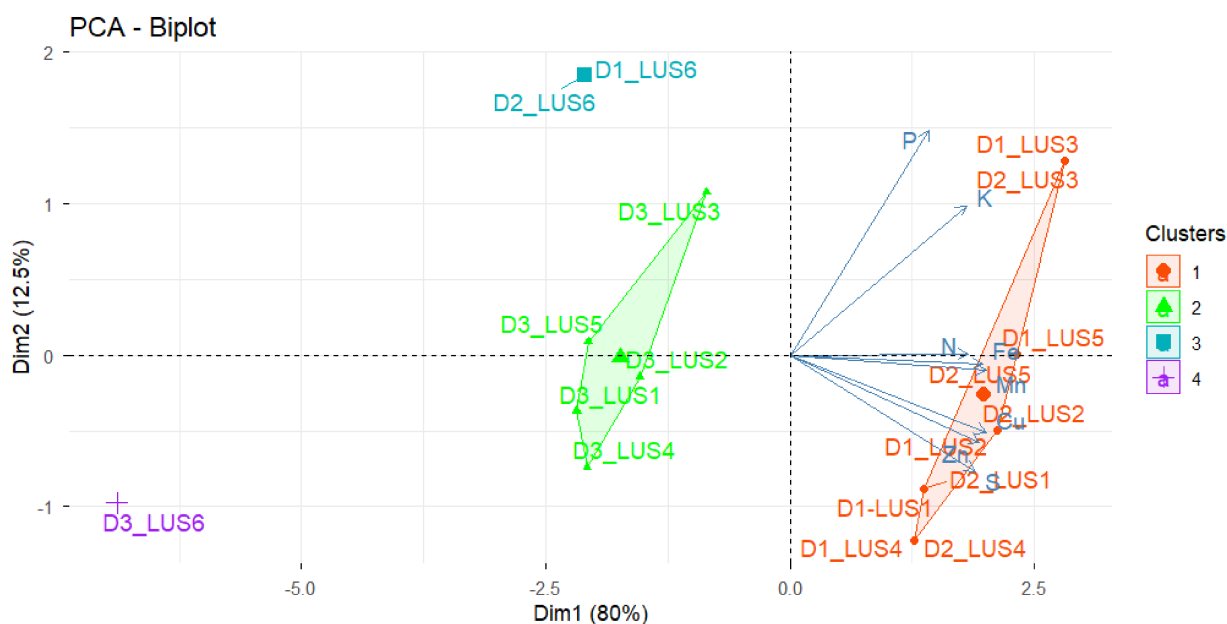
**Table 4.** Mean values of nutrient status different land use systems in the Central Plain Zone of Uttar Pradesh at different depths of soil samples.

LUS	Depth	N (Kg ha <sup>-1</sup> )	P (Kg ha <sup>-1</sup> )	K (Kg ha <sup>-1</sup> )	S (mg kg <sup>-1</sup> )	Fe (mg kg <sup>-1</sup> )	Mn (mg kg <sup>-1</sup> )	Cu (mg kg <sup>-1</sup> )	Zn (mg kg <sup>-1</sup> )
LUS1	D1	252.50 c ± 24.94	12.56 c ± 1.31	277.72 d ± 8.00	78.15 ab ± 5.38	164.12 a ± 5.35	52.42 b ± 6.58	2.61 a ± 0.59	1.45 b ± 0.15
	D2	252.50 c ± 25.11	12.56 c ± 1.28	277.72 d ± 8.41	78.15 ab ± 7.36	164.12 a ± 5.31	52.42 b ± 7.43	2.61 a ± 0.53	1.45 b ± 0.15
	D3	156.00 d ± 13.68	10.57 b ± 1.51	229.06 c ± 11.79	43.11 b ± 3.48	80.65 a ± 8.96	35.31 a ± 3.25	1.80 a ± 0.15	0.72 a ± 0.16
LUS2	D1	272.67 b ± 3.93	14.58 b ± 0.31	303.70 c ± 10.35	77.32 b ± 6.65	159.85 a ± 27.21	51.85 b ± 5.43	2.85 a ± 0.30	1.80 a ± 0.23
	D2	272.67 b ± 5.60	14.58 b ± 0.29	303.70 c ± 10.27	77.32 b ± 6.38	159.85 a ± 27.22	51.85 b ± 5.43	2.85 a ± 0.30	1.80 a ± 0.23
	D3	245.84 b ± 4.02	11.17 b ± 0.32	259.55 b ± 2.34	51.70 a ± 14.25	81.09 a ± 19.12	29.20 a ± 6.74	1.79 a ± 0.37	0.80 a ± 0.21
LUS3	D1	330.45 a ± 7.87	17.98 a ± 0.07	395.22 a ± 3.81	77.97 ab ± 5.48	164.11 a ± 5.38	52.37 b ± 6.56	2.61 a ± 0.57	1.42 b ± 0.14
	D2	330.45 a ± 19.99	17.98 a ± 0.10	395.22 a ± 31.73	77.97 ab ± 5.14	164.11 a ± 5.37	52.37 b ± 6.56	2.61 a ± 0.57	1.42 b ± 0.14
	D3	287.40 a ± 2.02	14.63 a ± 1.47	295.28 a ± 7.52	43.04 b ± 3.56	80.78 a ± 9.13	35.38 a ± 3.18	1.81 a ± 0.16	0.72 a ± 0.17
LUS4	D1	258.62 bc ± 9.37	11.10 d ± 0.53	277.72 d ± 8.00	78.15 ab ± 5.38	164.12 a ± 5.35	52.42 b ± 6.58	2.61 a ± 0.59	1.45 b ± 0.15
	D2	258.62 bc ± 9.00	11.10 d ± 0.40	277.72 d ± 13.46	78.15 ab ± 5.31	164.12 a ± 5.77	52.42 b ± 6.58	2.61 a ± 0.59	1.45 b ± 0.15
	D3	210.78 c ± 16.84	8.96 c ± 0.38	229.06 c ± 11.79	43.11 b ± 3.48	80.65 a ± 8.96	35.31 a ± 3.25	1.80 a ± 0.15	0.72 a ± 0.16
LUS5	D1	275.05 b ± 3.66	15.83 b ± 0.37	317.55 b ± 2.33	84.33 a ± 2.69	156.97 a ± 13.66	60.89 a ± 5.43	2.70 a ± 0.73	1.49 b ± 0.14
	D2	275.05 b ± 3.98	15.83 b ± 0.38	317.55 b ± 2.80	84.33 a ± 2.60	156.97 a ± 13.85	60.89 a ± 5.43	2.70 a ± 0.73	1.49 b ± 0.14
	D3	245.06 b ± 2.88	11.01 b ± 0.51	252.74 b ± 1.64	50.71 b ± 5.01	74.86 a ± 14.78	30.72 a ± 8.60	1.28 b ± 0.31	0.59 a ± 0.18
LUS6	D1	122.79 d ± 11.41	15.74 b ± 2.09	279.86 d ± 17.67	15.15 c ± 1.26	113.24 b ± 0.88	35.48 c ± 0.75	1.32 b ± 0.35	0.60 c ± 0.09
	D2	122.79 d ± 7.28	15.74 b ± 1.97	279.86 d ± 17.75	15.15 c ± 1.23	113.24 b ± 0.87	35.48 c ± 0.75	1.32 b ± 0.35	0.60 c ± 0.09
	D3	30.55 e ± 3.64	6.40 d ± 0.76	94.32 d ± 10.44	10.96 c ± 0.73	5.59 b ± 0.62	3.43 b ± 0.66	0.53 c ± 0.13	0.23 b ± 0.04

**Land Use Systems:** LUS1—fallow land use system, LUS2—crop cultivation-based land use system, LUS3—horticulture-based land use system, LUS4—forest-based land use system, LUS5—vegetable-based land use system, and LUS6—barren land use system. **Depths:** D1 (0–15 cm), D2 (15–30 cm), and D3 (30–60 cm). **Units:** Kg ha<sup>-1</sup>—kilograms per hectare, mg ha<sup>-1</sup>—milligrams per kg soil. Letters denote significant difference among land use systems ( $p < 0.05$ ).

### 3.7. Principal Component Analysis

Our PCA of soil nutrients explained 92.5% of the total variability, providing a clear and reliable picture of soil fertility patterns across different sites. Principal Component 1/Dim1 (80%) separates samples along a horizontal gradient: Cluster 1 (orange) on the right is associated with higher levels of N, P, K, and other essential macronutrients, indicating nutrient-rich soils, while Clusters 2, 3, and 4 on the left have lower nutrient levels, reflecting less fertile or contrasting soil conditions (Figure 8). This shows that PC1 primarily captures the overall fertility gradient.



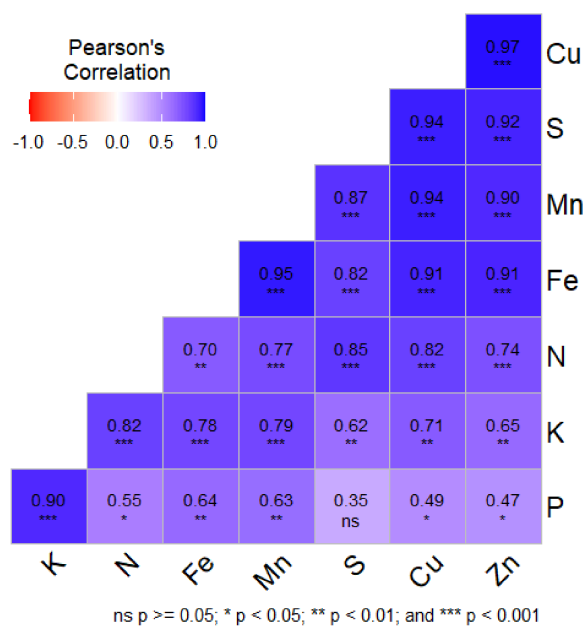
**Figure 8.** Principal component analysis (PCA) of soil nutrient status under different land use systems across varying soil depths. Land Use Systems: LUS1—fallow land use system, LUS2—crop cultivation-based land use system, LUS3—horticulture-based land use system, LUS4—forest-based land use system, LUS5—vegetable-based land use system, and LUS6—barren land use system. Depths: D1 (0–15 cm), D2 (15–30 cm), and D3 (30–60 cm).

Principal Component 2/Dim2 (12.5%) explains subtler differences along the vertical axis, highlighting nutrient balance and interactions, such as P levels and micronutrient dynamics. While it contributes less to total variance, PC2 provides valuable insight into how nutrients interact and vary independently of overall fertility. Together, PC1 and PC2 create a clear, two-dimensional snapshot of soil nutrient status, helping visualize both fertility levels and underlying nutrient relationships across the study sites.

### 3.8. Pearson's Correlation

Pearson's correlation analysis revealed strong and meaningful relationships among macro- and micronutrients, indicating coordinated nutrient dynamics within the soil system. Phosphorus (P) exhibited a very strong positive correlation with potassium (K) ( $r = 0.90$ ), suggesting shared sources, similar retention behavior, or synchronized nutrient release patterns. Nitrogen (N) showed moderate to strong associations with Fe, Mn, and Zn ( $r = 0.70$ – $0.77$ ), reflecting the influence of organic matter decomposition and microbial activity on both macro- and micronutrient availability (Figure 9). Among micronutrients, Fe, Mn, Cu, and Zn were highly intercorrelated ( $r = 0.82$ – $0.95$ ), highlighting common geochemical controls and possible co-precipitation or adsorption mechanisms. The strongest relationships were observed between Cu and S ( $r \approx 0.97$ ) and Mn and S and Cu ( $r = 0.90$ – $0.94$ ), emphasizing the role of redox processes and soil mineral interactions. Overall, these signif-

icant positive correlations indicate that improving soil fertility through balanced nutrient management and organic amendments may simultaneously enhance multiple nutrient pools, ultimately supporting sustainable crop productivity and soil health.



**Figure 9.** Pearson’s correlation analysis of soil nutrient status under different land use systems across varying soil depths.

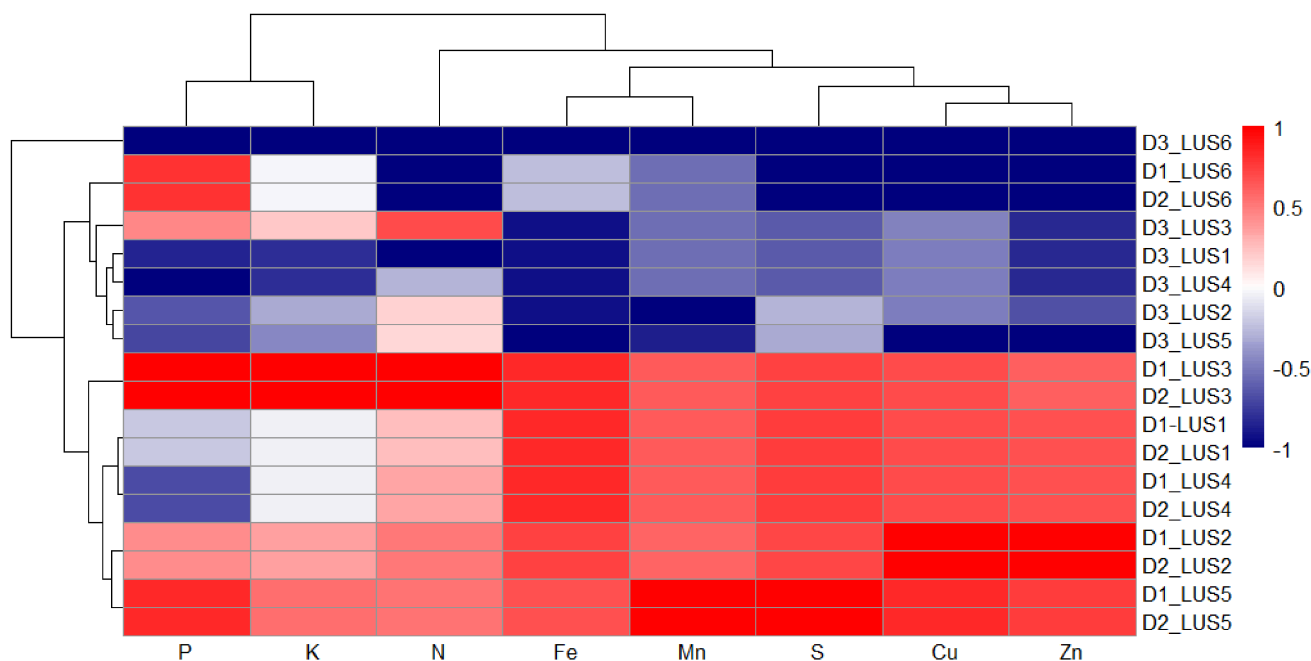
### 3.9. Heatmap Analysis

The hierarchical clustering heatmap demonstrates a clear vertical stratification of soil nutrients across LUS1–LUS6 and D1, D2, and D3, where D1 represents the surface layer, D2 the subsurface, and D3 the deeper soil horizon (Figure 10). Across all land use systems (LUS1–LUS6), samples primarily cluster according to depth rather than land use category, indicating that soil depth is the dominant factor influencing nutrient distribution. The D1 and D2 layers consistently group together and are characterized by higher standardized intensities of Fe, Mn, S, Cu, and Zn, along with comparatively elevated N, reflecting nutrient enrichment in upper soil horizons. This enrichment likely results from greater organic matter accumulation, litter deposition, root proliferation, and intensified microbial activity near the surface. In contrast, D3 samples from LUS1–LUS6 form a distinct cluster with predominantly lower standardized values, indicating reduced micronutrient availability and overall nutrient depletion in deeper layers. The nutrient dendrogram further shows coherent grouping of Fe–Mn and Cu–Zn, suggesting similar geochemical behavior and mobility patterns, while P and K exhibit relatively coordinated distribution among macronutrients. Overall, the heatmap highlights a pronounced depth-dependent nutrient gradient across LUS1–LUS6, with decreasing micronutrient intensity and altered macronutrient distribution from D1 to D3, underscoring the strong influence of soil depth on nutrient dynamics irrespective of land use system.

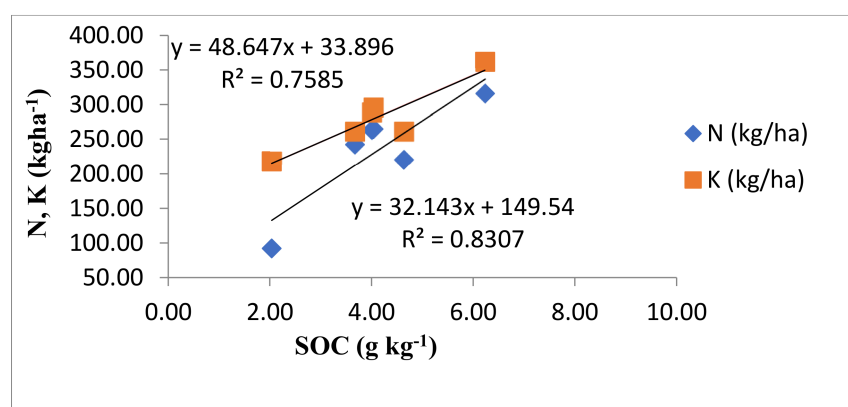
### 3.10. Analysis of Correlation and Linear Regression

Soil organic carbon (SOC) and the availability of key nutrients, such as nitrogen (N), phosphorus (P), potassium (K), sulfur (S), and micronutrients, such as zinc (Zn), copper (Cu), manganese (Mn), and iron (Fe), showed a strong positive linear relationship in all six LUSs (Figure 11). Zinc had the strongest correlation ( $R^2 = 0.9964$ ), followed by nitrogen ( $R^2 = 0.8307$ ), potassium ( $R^2 = 0.7585$ ), iron ( $R^2 = 0.5881$ ), manganese ( $R^2 = 0.5062$ ), sulfur ( $R^2 = 0.4999$ ), phosphorus ( $R^2 = 0.3903$ ), and copper ( $R^2 = 0.314$ ) in terms of their association

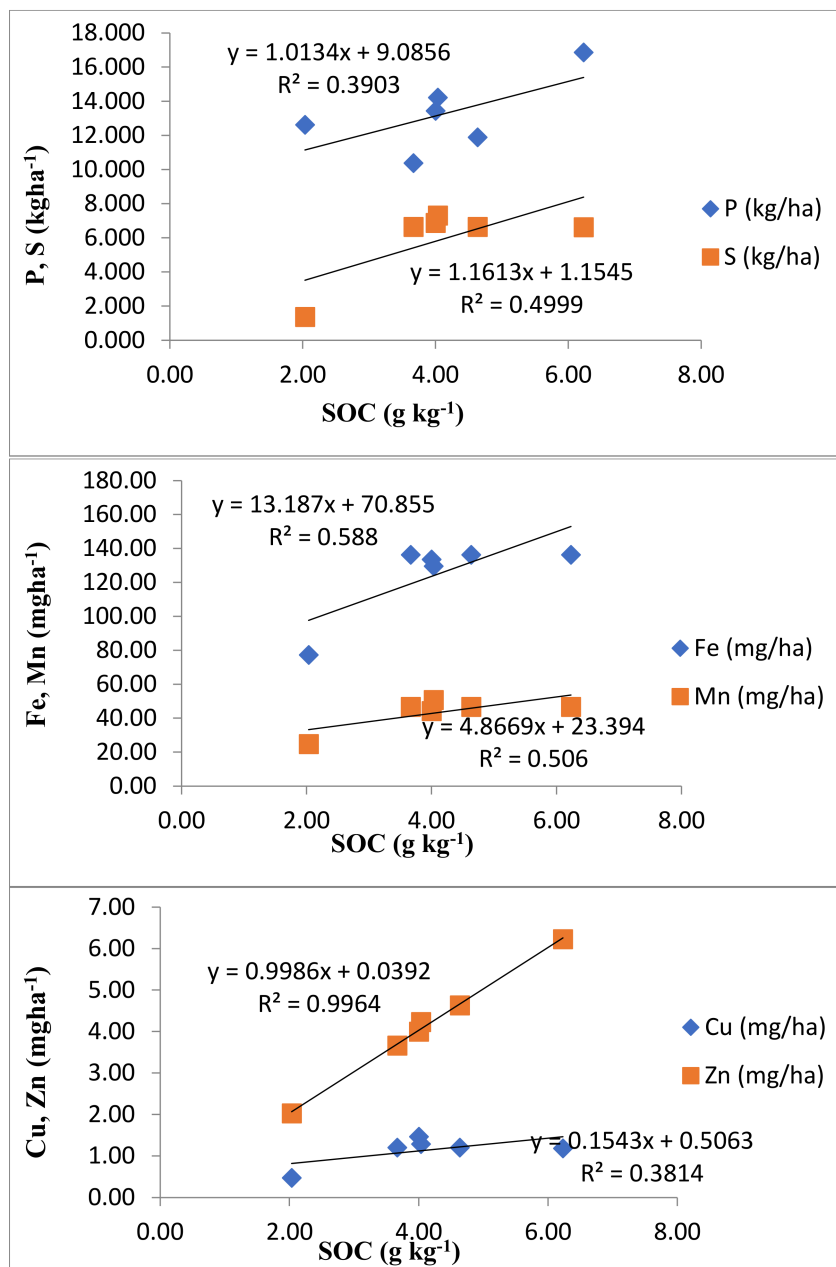
with SOC across different LUSs, according to the regression analysis. This implies that zinc concentration had the strongest effect on SOC, demonstrating its critical function in the accumulation of soil organic carbon [70]. The general pattern shows that the concentration of all the soil nutrients under investigation rises in tandem with the enrichment of SOC under various land use systems [64]. This highlights how land use practices and soil nutrient dynamics interact and how nutrient management techniques can increase the amount of organic carbon in the soil.



**Figure 10.** Heatmap analysis of soil nutrient status under different land use systems across varying soil depths. Land Use Systems: LUS1—fallow land use system, LUS2—crop cultivation-based land use system, LUS3—horticulture-based land use system, LUS4—forest-based land use system, LUS5—vegetable-based land use system, and LUS6—barren land use system. Depths: D1 (0–15 cm), D2 (15–30 cm), and D3 (30–60 cm).



**Figure 11.** Cont.



**Figure 11.** Relationship between mean value of SOC and available N, P, K, S, and micronutrients (Zn, Cu, Mn, and Fe) under different LUSs at 0–60 cm soil depth.

The results indicate a strong positive relationship between soil organic carbon (SOC) and the availability of both macro- and micronutrients across all land use systems. This suggests that higher SOC levels enhance nutrient availability in soils. Among all nutrients, zinc showed the strongest association with SOC, highlighting its key role in carbon accumulation, while copper showed the weakest relationship. Overall, the findings demonstrate that improved soil organic carbon through better land use and management practices can significantly enhance soil fertility and nutrient dynamics.

#### 4. Discussion

##### 4.1. Variation in Physico-Chemical Properties of Soils Across Land Use Types

The results clearly demonstrate that land use systems exert a strong influence on soil physico-chemical properties across the Central Plain Region of Uttar Pradesh. Soil reaction (pH) exhibited comparatively low variability among the studied parameters, yet distinct

differences were observed among land use systems. Lower pH under horticulture systems indicates mild acidification due to continuous organic inputs [71,72], whereas higher pH in barren lands reflects alkaline conditions driven by low biological activity and accumulation of basic cations [27,73]. The increase in pH with depth aligns with carbonate accumulation trends typical of alluvial and semi-arid soils [74].

The lowest EC values were typically found in relatively undisturbed and well-vegetated regions where minimal salt accumulation occurs due to efficient nutrient cycling [75], and the highest EC was observed in barren land systems, particularly in poorly managed or arid regions, where salt accumulation is more pronounced due to evaporation and lack of leaching. Among cultivated regions, crop-based systems showed relatively higher EC compared to other managed systems, followed by forest, fallow, and vegetable systems. This pattern reflects variations in fertilizer use, irrigation practices, and salt dynamics across regions. The increase in EC with soil depth across all land uses indicates downward movement and accumulation of soluble salts in subsoil layers, which is consistent with processes of leaching and ion translocation in alluvial soils [76].

The highest SOC content was observed in horticulture-based systems, particularly in regions with dense perennial vegetation and regular organic matter addition through litter fall and minimal soil disturbance [44,54,77]. In contrast, the lowest SOC was recorded in barren land systems, especially in degraded and sparsely vegetated regions where organic inputs are minimal, and decomposition rates may exceed accumulation [78]. Forest-based systems also maintained relatively higher SOC compared to cultivated lands, owing to continuous biomass addition and stable ecological conditions. Conversely, crop-based and fallow systems exhibited comparatively lower SOC levels, likely due to intensive cultivation, residue removal, and insufficient organic amendments. These regional differences emphasize the importance of sustainable land management practices in maintaining soil carbon and overall soil health [79].

Bulk density (BD) increased with depth, with lower values in surface soils of horticulture and forest systems due to higher organic matter and better aggregation [80]. Higher BD in mid-depth layers of cultivated and barren lands indicates compaction effects from low organic inputs and mechanical disturbance. Soil texture showed increasing silt and clay with depth, indicating illuviation processes influencing water retention and nutrient availability [51].

#### *4.2. Soil Organic Carbon Fractions and Stock Dynamics Under Various Land Use Systems*

The observed variations in soil organic carbon (SOC) fractions and associated biological properties across land use systems are well supported by region-specific ecological and management factors. The higher accumulation of very labile, labile, less labile, and non-labile carbon fractions in LUS3 can be justified by the presence of dense vegetation cover, continuous litter deposition, and higher root biomass, which collectively enhance organic matter input and carbon stabilization [81]. In such systems, minimal soil disturbance and sustained organic additions promote aggregation and protection of carbon within soil microstructures, leading to overall improvement in SOC pools. In contrast, the lower values of carbon fractions in LUS6 are likely due to intensive cultivation, reduced biomass return, and frequent soil disturbance, which accelerate organic matter decomposition and carbon loss [82].

The lack of protective vegetation cover further exposes soil to erosion and oxidation processes, resulting in depletion of both labile and stable carbon pools [83]. These conditions are typical of degraded or poorly managed land use systems in the region. The relatively higher microbial biomass carbon (MBC) and dehydrogenase activity (DEA) observed in LUS4 and LUS5 can be attributed to favorable soil microclimatic conditions, including better

moisture retention, moderate temperature, and higher availability of easily decomposable organic substrates. These environments support active microbial communities, which in turn enhance enzymatic activities and nutrient cycling [84].

On the other hand, the reduced microbial activity in LUS6 is justified by limited substrate availability and unfavorable soil conditions, which restrict microbial growth and function. The decline in SOC fractions and biological activity with increasing soil depth is primarily due to reduced organic matter input from plant residues and lower root density in subsurface layers [85]. Additionally, limited aeration and slower microbial processes at deeper depths contribute to reduced carbon turnover [86]. However, the persistence of less labile and non-labile carbon fractions in deeper layers is justified by their recalcitrant nature and protection within soil aggregates, making them less susceptible to decomposition. Furthermore, variations in hot water-soluble carbon (HWSC) and acid-hydrolyzable humic carbon (AHHC) across land use systems can be explained by differences in organic matter quality and decomposition rates. Systems with higher organic inputs and active microbial processes tend to have greater soluble and hydrolyzable carbon fractions, indicating improved soil functional stability [87].

#### *4.3. Soil Nutrient Dynamics Under Different LUSs*

The results clearly demonstrate that nutrient dynamics are strongly governed by land use systems (LUSs), with clear justification for both the increase and the decrease in nutrient availability across systems and soil depths. Horticultural land use system soils recorded higher available nitrogen (N) compared to crop, vegetable, forestry, fallow, and barren systems [88]. This increase is mainly due to continuous addition of organic residues through leaf litter fall and root biomass deposition, which enhances microbial activity and accelerates nitrogen mineralization. In contrast, agricultural lands under intensive tillage show reduced N availability because frequent soil disturbance promotes rapid oxidation of soil organic matter, thereby lowering the nitrogen pool. The lowest N in barren land is due to minimal vegetation cover, which results in poor organic matter input and higher leaching losses through rainfall and irrigation. The decline in N with soil depth is primarily attributed to reduced organic carbon content and increased immobilization of nitrogen in deeper layers, where microbial activity is also limited [89].

Similarly, higher phosphorus (P) availability under horticultural systems is justified by greater organic matter content, which facilitates the release of organically bound phosphorus through mineralization [90]. In addition, continuous application of phosphatic fertilizers in managed systems contributes to its accumulation. Lower P in other systems, particularly barren land, is due to a lack of organic inputs and limited biological activity, which restricts phosphorus cycling [91].

Potassium (K) availability is relatively higher in horticulture- and forest-based systems due to the continuous addition of organic matter in the form of litter and root residues. These inputs enhance the weathering of minerals and solubilization of otherwise unavailable forms of potassium. In contrast, reduced K in barren land is associated with soil degradation and leaching losses. Moreover, in soils with higher salinity or alkalinity, nutrient imbalance and ionic interference restrict potassium uptake by plants, further contributing to its lower availability [92].

Micronutrient (Fe, Mn, Cu, Zn) availability also varied significantly across LUSs and depths, largely influenced by organic matter content. Higher micronutrient concentrations in surface soils of fallow, vegetable, and cultivated systems are due to greater organic inputs, which enhance chelation and solubility of these nutrients. Increased microbial activity in these systems further promotes nutrient transformation and availability [93]. On the other hand, barren land, especially in subsurface layers, showed the lowest micronutrient

content due to poor organic matter status, weak biological activity, and limited nutrient recycling [94]. The observed decrease in micronutrients with soil depth is again linked to declining organic matter, reduced root activity, and weaker microbial processes. The relatively higher Fe and Mn status in the region suggests that these elements are inherently sufficient in tropical soils, but their availability is still closely regulated by organic matter and soil biological activity [95]. Although the studied soils were predominantly alluvial and contained negligible visible coarse fragments, the coarse fragment fraction (>2 mm) was not quantified separately. Therefore, SOC stock estimations should be interpreted cautiously in soils where gravel or stone content may be appreciable. Overall, perennial and diversified land use systems improved soil carbon dynamics, microbial activity, and nutrient availability, highlighting their importance for sustainable land management and climate-resilient agroecosystems.

## 5. Conclusions

Land use change influences a number of biological and physiological processes. Poor land use decisions lead to land degradation and reduce soil health. Results indicated that soil pH and electrical conductivity increased with depth, while organic carbon decreased, reflecting higher surface organic matter from litter and microbial activity. Bulk density was highest in the middle layer, and soil moisture remained relatively stable. Soil texture varied with depth, with decreasing sand and increasing silt and clay. SOC fraction analysis covering very labile, labile, less labile, and non-labile carbon, along with microbial parameters (HWSC, AHHC, MBC, and DEA), revealed that labile carbon is particularly sensitive to land use. Among the six land use systems, horticulture-based systems (LUS3) consistently exhibited the highest SOC fractions, microbial activity, and nutrient status, whereas barren lands (LUS6) showed the lowest levels. SOC fractions and nutrient reserves declined with depth, emphasizing the importance of topsoil in carbon cycling. Barren lands exhibited the most alkaline pH, the highest salinity risk, and the poorest organic and nutrient content. In contrast, horticulture-based systems maintained favorable soil conditions, including lower pH, higher SOC, moderate electrical conductivity, and the highest surface reserves of macronutrients (N, P, K, S) and micronutrients (Fe, Mn, Cu, Zn). This highlights the crucial role of continuous organic inputs, such as leaf litter and root biomass, in enhancing soil fertility. The findings underscore the need for sustainable land management practices that incorporate organic residues, minimize soil disturbance, and employ targeted interventions such as organic amendments, salt-leaching, and pH adjustment to restore fertility in degraded or barren soils. Future research should expand spatially and temporally to refine land use guidelines that optimize crop productivity, enhance carbon sequestration, and ensure long-term soil resilience.

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**Data Availability Statement:** The datasets used and analyzed during the current study are available from the corresponding authors upon reasonable request.

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